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EXECUTIVE SUMMARY

This report contains the results of the first multi-annual geospatial, geochemical and oceanographic study of the stranding of coastal bitumen (asphaltite and tar balls) on representative ocean beaches (n = 31) along the entire coastline of South Australia. Previous investigations of this phenomenon, dating back to the mid-1800s, were episodic and limited to beaches on the Limestone Coast, southern Kangaroo Island and the foot of Eyre Peninsula. The present study revisited the same 31 beaches after winter in 2014, 2015 and 2016, collecting a total of 631 specimens, and in so doing filled a major gap in our knowledge of bitumen stranding along the western side of Eyre Peninsula to the head of the Great Australian Bight.

This investigation employed an updated version of the NOAA SCAT method for assessing shoreline oiling conditions in order to establish the natural baseline hydrocarbon loading of the South Australian coastline. Major findings from this part of the project are:

- Tar balls (waxy bitumen) preferentially strand in the upper intertidal to upper shore zones of southwest-facing ocean beaches, whereas the less common, denser asphaltites tend to accumulate on beaches with a northwest aspect.
- The number and size of individual strandings is considerably less than reported for the period 1960-1995.
- Waxy bitumens strand along the entire South Australian shoreline, with a particular focus on the Limestone Coast.
- Asphaltites are more commonly found along the west coast of Eyre Peninsula, suggesting different transport mechanisms and/or a different point of origin.
- The high abundance of waxy bitumens with geochemical similarities to Indonesian-sourced oils is consistent with the transport of these materials as flotsam, first in the southward Leeuwin Current which then feeds into the eastward-flowing coastal and South Australian currents.

The resulting catalogued collection of asphaltites and tar balls, the largest yet assembled for any geochemical investigation of this type in Australia, underwent detailed physical, elemental, isotopic and biomarker characterisation. The principal findings of this work are:

- Fifteen oil families and sub-families of South Australian coastal bitumen were identified, significantly more than the six families known from previous studies.
- The waxy bitumens differ from those found along the same coastline between 1960 and 1995 in being much more weathered and biodegraded, consistent with both their smaller size and lesser numbers.
- Soft asphaltic bitumen recovered from a site on the Limestone Coast is unique, and likely represents an early expulsion product from a Cretaceous marine source rock similar to that which gave rise to the asphaltites. Both point to the likely existence of an active petroleum system in the Bight Basin.
- The majority of the waxy bitumens have distinctive Cenozoic lacustrine biomarker signatures which mark their origin from distant offshore oil seeps in the Indonesian Archipelago.
- The abundance of both asphaltite and waxy bitumen stranding on South Australian beaches has declined dramatically over the past 20+ years, attesting to a diminution of seep activity.
and/or, in the case of the latter, improved environmental practices related to tanker washing and oil spillage within Indonesian waters.

- At least two waxy bitumen sub-families lack Indonesian signatures and could therefore originate from seeps in offshore basins along Australia’s western and southern continental margin.
- The discovery of several new oil families of non-Indonesian provenance among the ocean wanderer bitumens that impact the South Australian coastline has potentially important implications for the prospectivity of its offshore sedimentary basins.

For the first time oceanographic modelling was employed to develop an understanding of the transport of coastal bitumens in the Great Australian Bight (GAB). The findings show that:

- Oceanographic models which simulate currents alone do not adequately account for the observed stranding locations of tarballs and asphaltite.
- Models which simulate both currents and Stokes drift can describe their stranding locations across the region.
- Winter strandings are dominated by materials transported from west to east by the Leeuwin Current and storms.
- The winter modelled conditions can supply materials to all of the beaches in the GAB from areas along the shelf break at the western extent of the model, consistent with the discovery of Indonesian-sourced waxy bitumens on all the beaches surveyed.
- The winter modelling shows that there is a low probability of bitumens sourced in the Ceduna or Duntroon sub-basins reaching the coastline at the Head of the Bight and on the northern Eyre Peninsula.
- Summer strandings are dominated by materials transported from the southeast by Stokes drift caused by strong northerly to northwesterly winds.
- Throughout all seasons there is a low probability of any materials sourced in the Morum sub-basin travelling to the west of Kangaroo Island and encountering the ocean beaches on Eyre Peninsula.
- The area with the highest frequency of particle tracks is located to the west of Kangaroo Island and overlies the Duntroon sub-basin and the eastern part of the Ceduna sub-basin.
- Any locally derived materials within this area can reach all beaches in the study area and there is seismic support for potential leakage indicators in this area.

These research findings have addressed and exceeded the original project objectives and offer an unprecedented insight into the transport mechanisms, stranding processes, alteration and source of coastal bitumens which are stranded along the coastlines of South Australia. The data produced forms an important, and continuing, baseline for hydrocarbon loadings on the beaches of South Australia and whilst a definitive origin of the new tarball oil families and asphaltites has not been established they provide evidence of active hitherto unknown regional petroleum systems.
Part I  Introduction
INTRODUCTION

Overview
BP Developments Australia, CSIRO, the South Australian Research and Development Institute (SARDI), the University of Adelaide and Flinders University are working on a 4 year, $20 million research program to improve the understanding of the environmental, economic and social value of the Great Australian Bight (GAB). The GAB Research Program is administered by a Management Committee with representation from BP, CSIRO, SARDI, Adelaide University and Flinders University and is advised by an independent Science Panel (ISP) comprising internationally recognised experts in key discipline areas.

The GAB Research Program comprises 7 themes and 16 research projects. This report presents the findings of research conducted for Theme 5 (Petroleum Geology & Geochemistry) and Project 5.2 (Asphaltite and tarball surveys).

Background and need

The Great Australian Bight (GAB), and particularly the Ceduna Sub-basin, is considered one of the most prospective deepwater frontier basins in offshore Australia (Totterdell et al., 2008). It contains up to 15 km of mid to late Cretaceous deltaic and marine sediments within the Tiger and Hammerhead Supersequences that provide potential reservoirs, seals and oil-prone source rocks at several stratigraphic levels (Blevin et al., 2000; Totterdell et al., 2000; Struckmeyer et al., 2001).

At the start of the Great Australian Bight Research Program (GABRP) no liquid hydrocarbons had been directly measured in the GAB and therefore no active petroleum system was proven. This was despite the determination of possible hydrocarbon migration routes and accumulation zones within some of the few exploration wells drilled in the basin (i.e. the detection of oil inclusions at low abundance) and indications that there are hydrocarbons naturally leaking into the GAB (i.e. asphaltite strandings).

Previous geochemical studies of the occurrence of coastal bitumens (McKirdy et al., 1986, 1994; Currie et al., 1992; Padley et al., 1993; Padley, 1995) have shown that the vast majority of the tarballs found are waxy bitumens which originated in the Indonesian Archipelago and were subsequently carried into South Australian waters by the South Equatorial and Leeuwin Currents. Notwithstanding their extra-GAB provenance, they are the dominant contributor to the natural hydrocarbon loading of southern Australia’s coastline (Edwards et al., 2016). Moreover, as surface drifters (specific gravity 13–40° API) their stranding pattern provides a predictive template for the ultimate destination of any oil spill that might arise from exploration or production operations in the GAB.

The asphaltites are almost certainly of local origin, although the exact location of their parent petroleum system has yet to be established (Edwards et al., 1998, 1999; Totterdell et al., 2008; Hall et al., 2014). Their quasi-neutral buoyancy (4–18° API) implies that they were submerged or even bottom drifters. Hence, it is not surprising that their historical stranding pattern should differ markedly from that of the pelagic tarballs (see Padley, 1995; Edwards et al., 2016).
A detailed inventory of the different varieties of petroleum hydrocarbons (oil slicks, tarballs, asphaltites) known to strand along the adjacent South Australian coastline is required in order to 1) identify those most likely to originate from offshore hydrocarbon seeps, and 2) provide a baseline understanding of the nature and abundance of asphaltites and tarballs on the coastline before any major exploration or production of oil/gas commences in the Bight.

Objectives
In order to better constrain the provenance of coastal bitumen in the Great Australian Bight (GAB) region, Project 5.2 of the GABRP aimed to address key knowledge gaps identified within the literature:

1. The contemporary natural hydrocarbon loading of the South Australian coastline.

2. The sites of most frequent present-day (as opposed to historical) stranding of asphaltite and tarballs, and their geographic relationship to offshore sea surface currents.

3. Molecular and isotopic characteristics of a statistically significant population of freshly stranded asphaltites, necessary to determine their degree of weathering and hence their likely proximity to the parent seep(s).

To address these knowledge gaps Project 5.2 objectives were to:

1. Map and quantify the contemporary natural hydrocarbon loading of selected ocean beaches on Eyre Peninsula, Kangaroo Island and the Limestone Coast. The latter two areas being known sites for the regular stranding of asphaltites, waxy bitumens (tarballs) and/or oil slicks. The collection of tarballs and asphaltites on the beaches of South Australia will allow their geochemistry to be documented, and ascertain their probable origins using oceanographic and geological models.

Establishing the present-day hydrocarbon loading of the South Australian coastline (its geographic distribution, concentration in kg/km, compositional heterogeneity and principal sources) will act as a baseline against which the environmental impact of any spillage arising from future drilling operations in the GAB can be measured, and also serve as a predictive template for the likely destination(s) of such accidental spillage.

2. Identify the provenance and weathering of the asphaltites and any non-Indonesian tarballs and oil slicks using their elemental, isotopic and molecular fingerprints. In particular, their source and maturity-specific parameters will provide essential clues to their origin. These data, in combination with stranding distributions (GABRP Project 5.2 Objective 1); oceanographic modelling using a detailed knowledge of both the sea surface currents traversing the continental margin of South Australia, including zones of upwelling (Oceanography Theme); geological interpretations of leakage indicators and offshore seepage studies (Seeps and Leakage GABRP Project 5.1); and other petroleum migration, timing and accumulation indicators (Fluid Inclusions GABRP Project 5.3), will help delineate possible seafloor hydrocarbon leakage points within the GAB.
REGIONAL GEOLOGY OF THE BIGHT BASIN

The Bight Basin formed during the break-up of eastern Gondwana in the Late Jurassic–Early Cretaceous. It extends for ~2000 km along the Australian southern margin and comprises a series of extensional depocentres in modern day water depths between 200 m and >4000 m. No significant hydrocarbons have been found in the basin, which remains an exploration frontier.

Basin Outline

The Bight Basin is a large, mainly offshore basin situated along the western and central parts of the continental margin of southern Australia (Figure 1). The basin extends from the Leeuwin Fracture Zone in the west, to just south of Kangaroo Island in the east, where it adjoins the Otway Basin. The basin contains a number of depocentres that were formed by the rifting and thinning of the Australian Plate; including the Ceduna, Duntroon, Eyre, Bremer and Recherche sub-basins (Figure 1). The current sub-divisions of the Bight Basin are based on those defined by Bradshaw et al., (2003) and Totterdell & Bradshaw (2004) and have evolved from many previous basin definitions published by various researchers. Within the revised definition of the Bight Basin, the former Duntroon Basin has been remapped as a smaller, genetically related rift system along the margin of the main rift basin, and has been reclassified as a sub-basin. The Bight Basin is overlain unconformably by dominantly cool-water carbonates of the Cenozoic Eucla Basin (Messent, 1996; Feary & James, 1998).

Figure 1: Location of the Bight Basin with component sub-basins. © Commonwealth of Australia (Geoscience Australia) 2016.
Basin evolution and tectonostratigraphic framework

The Bight Basin is a large Mesozoic to Cenozoic depocentre, which formed during the breakup of Gondwana (Fraser & Tilbury, 1979; Bein & Taylor, 1981; Willcox & Stagg, 1990; Hill, 1995; Totterdell et al. 2000; Norvick & Smith, 2011; Teasdale et al., 2003). The basin developed along Australia’s southern margin during a period of extension and passive margin evolution that commenced in the Middle–Late Jurassic (Teasdale, 2003). The Ceduna Sub-basin, which is the focus of current hydrocarbon exploration activities, contains in excess of 15 km of syn- and post-riift Mesozoic sediments (Figure 2-4). It occurs in water depth ranging from 200 m to over 4000 m and has an area of approximately 90,000 km² (Sommerville, 2001). The northern margin is characterised by a series of fault-bound half grabens that contain Middle Jurassic to Early Cretaceous syn-riift fill (Figure 4). The southwestern boundary is interpreted at the basinward edge of an associated toe-thrust zone. The sub-basin is characterised by five main phases of evolution (King & Mee, 2004; Totterdell & Bradshaw, 2004; Blevin & Cathro, 2008):

- A Late Jurassic Early Cretaceous mechanical subsidence phase due to a phase of intracontinental extension. Extensional deformation in the Ceduna Sub-basin appears to have been focused along a pre-existing NW–SE-trending margin of the Gawler Craton resulted in simple half graben along the eastern margin of the sub-basin. The rift fill comprises the Callovian–Kimmeridgian Sea Lion Supersequence and the Tithonian–early Berriasian Minke Supersequence (Figure 5).

- An Early Cretaceous phase of slow thermal subsidence represented by the largely non-marine Berriasian Southern Right Supersequence and the Valanginian to mid-Albian Bronze Whaler Supersequence (Figure 5). The onlapping, sag-fill geometry of the succession suggests that accommodation was created largely by thermal subsidence and compaction, with deposition concentrated over the earlier half graben.

- A second period of active extension during the Early–Late Cretaceous causing rapid subsidence in the sub-basin. The high subsidence rates recorded from the mid-Albian were possibly controlled by an interplay between thermal subsidence, mechanical extension and gravity-driven growth faulting (Mulgara Fault Family) in the White Pointer delta. This period of accelerated subsidence continued until the commencement of sea-floor spreading between Australia and Antarctica in the Late Santonian, and coincided with a period of rising global sea level. This resulted in a major marine flooding event, and widespread deposition of the marine Blue Whale Supersequence, followed by deposition of the White Pointer and Tiger supersequences (Figure 5). Gravity-driven, detached extensional and contractional faults formed during the Cenomanian as a result of deltaic progradation. The deposition of the Tiger Supersequence coincides with a period of upper crustal extension resulting in the formation of large displacement, almost west–east-striking faults and the reactivation of many Cenomanian growth faults.

- The Late Cretaceous break-up associated with the first true oceanic crust at ~83 Ma (Sayers et al., 2001). This coincides with the base of the Hammerhead Supersequence (Figure 5). This boundary can be strongly erosional and characterised by significant incisions believed to be the result of uplift related to the commencement of sea-floor spreading.

- A post break-up thermal subsidence phase that coincided with the second, large progradational delta development resulting from a massive influx of sediments giving rise to the characteristic, prograding shelf-margin geometries of the Ceduna Delta succession with
the deposition of the Hammerhead Supersequence (Figure 5). A localised region of gravity-driven structures and simple, planar normal faulting developed with the latter reactivating older faults. Faulting appears to be latest Maastrichtian–Early Paleocene in age. Totterdell and Bradshaw (2004) suggest that the faulting was probably related to flexure of the margin caused by sediment loading during the Late Cretaceous.

Half-spreading rates were extremely slow from the Campanian (~83 Ma) until the Middle Eocene (~43 Ma), reaching a maximum of around 10 mm/year, although spreading rates were generally much less. From around 43 Ma, half-spreading rates rapidly increased to about 20 mm/year in a N-S orientation (Tikku & Cande, 1999). In the post-Eocene period, deposition of marine carbonates reflects deepening water and the end of the effect of regional tectonics on the development of the Bight Basin (Fraser and Tilbury, 1979).

Late Cretaceous and Tertiary igneous rocks have been interpreted on seismic data (Totterdell et al., 2000) and Tertiary volcanic rocks have been dredged from the Bight Basin (Davies et al., 1989; Clarke & Alley, 1993). Schofield & Totterdell (2008) detailed the distribution of the volcanic and intrusive bodies in the Ceduna Sub-basin and related them to the acceleration in seafloor spreading during the Middle Eocene.

Figure 2: Cross-section through the eastern Madura Shelf and Ceduna Sub-basin. From Bradshaw et al., 2003.
Figure 3: Cross-section through the eastern Madura Shelf, Ceduna Sub-basin and eastern Recherche Sub-basin. From Bradshaw et al., 2003.

Figure 4: Cross-section through the eastern Madura Shelf, Ceduna Sub-basin and eastern Recherche Sub-basin. From Bradshaw et al., 2003.
Figure 5: Bight Basin stratigraphic correlation chart showing basin phases and predicted source rock intervals (modified from Blevin et al., 2000 and Totterdell et al., 2000). The sea level curve (Haq et al., 1988) is modified to the time scale of Gradstein et al., 2004.
Exploration history

Petroleum exploration in the eastern Bight Basin (Eyre, Ceduna and Duntroon sub-basins) has occurred in three major cycles – the late 1960s to early 1970s, the early 1990s, and 2000–present (see O’Neil, 2003). After nearly fifty years of exploration in the offshore Bight Basin, only fourteen petroleum exploration wells have been drilled; Apollo-1 (1975), Borda-1 (1993), Columbia-1 (1982), Duntroon-1 (1986), Echidna-1 (1972), Gemini-1/1A (1975), Gnarlyknots-1/1A (2003), Greenly-1 (1993), Jerboa-1 (1980), Mercury-1 (1981), Platypus-1 (1972), Potoroo-1 (1975) and Vivinne-1 (1993). With the exception of Gnarlyknots-1/1A, all wells have been drilled in relatively shallow water near the basin margin. No significant hydrocarbons have been discovered and the deep-water parts of the basin remain largely untested and a frontier area.

During the last phase of exploration in the 2000s, exploration permits were awarded to a joint venture operated by Woodside Energy in the Ceduna Sub-basin. The joint venture acquired approximately 15,400 line km of 2D seismic data and in 2003, drilled Gnarlyknots-1/1A (Tapley et al., 2005). The well was the first and only attempt to test the deep-water Ceduna Sub-basin and failed to reach its primary target objective due to adverse weather conditions.

In 2007, Geoscience Australia embarked on a regional geological and sampling survey of the Bight Basin as part of the Australian Government’s Offshore Energy Security program. The survey targeted and sampled potential source rocks of late Cenomanian to early Turonian age on the northwest margin of the Ceduna Sub-basin. Analytical results indicated that the organic-rich rocks recovered by the survey are capable of generating liquid hydrocarbons (Totterdell et al, 2008; Totterdell & Mitchell, 2009). In 2009, six areas in the central Ceduna Sub-basin were made available for bidding and, in January 2011, BP Developments Australia Pty Ltd was awarded four permits (EPP 37–40).

Petroleum geology

The eastern part of the Bight Basin (Ceduna, Eyre and Duntroon sub-basins) is potentially one of the most prospective deepwater frontier basins in offshore Australia. The thick sedimentary succession, in excess of 15 km thick in places, and its evolution from local half grabens depocentres during the Jurassic, to an extensive sag basin in the Early Cretaceous and passive margin during the Late Cretaceous to Holocene, implies that there is significant potential for the presence of multiple petroleum systems.

The most prospective petroleum systems are believed to be associated with thick mid to Late Cretaceous deltaic and marine sediments (Blue Whale, White Pointer, Tiger and Hammerhead supersequences), which provide reservoirs, seals and potential oil-prone source rocks at several stratigraphic levels (Figure 5), together with a wide range of structural and stratigraphic plays (Blevin et al., 2000; Totterdell et al., 2000; Struckmeyer et al., 2001; Totterdell et al., 2008) – Figure 5.

Despite mounting evidence for oil-prone marine source rocks in the Ceduna Sub-basin (Totterdell et al., 2008), and speculation they might be present in distal facies of deep water areas in the Ceduna Sub-basin, hydrocarbon generation is a key uncertainty for explorers.
Source rocks

The key interpreted source rock intervals in the Bight Basin were deposited during a transgressive half-cycle which lasted from 165 to 84 Ma and was principally controlled by the propagation of the Southern Rift System from west to east (Stagg et al., 1990). Transgressive sequences began with non-marine facies of Callovian to Berriasian lacustrine shale, grading upwards into Valanginian to mid-Albian coastal plain and marginal marine facies composed of coal, sandstone and siltstone, and finally into fully marine facies of mid-Albian to late Santonian sandstone, siltstone and shale (Blevin et al., 2000). Therefore, the oldest potential source rocks in the Bight Basin are interpreted to be lacustrine and other non-marine facies within the Upper Jurassic to lowermost Cretaceous syn-rift succession of the Sea Lion and Minke Supersequences (Geoscience Australia, 2010). Indeed, algal-rich lacustrine shales from these supersequences are interpreted as the source for a breached palaeo-oil accumulation identified from oil stains in the Jerboa-1 well in the Eyre Sub-basin (Totterdell et al., 2000; Ruble et al., 2001).

In the overlying Southern Right Supersequence, fluvio-lacustrine claystone, siltstone and sandstone has been noted and interpreted to reach a maximum thickness of 1500 m in the Ceduna Sub-basin (Geoscience Australia, 2010). However, based on sparse Total Organic Carbon (TOC) and Rock-Eval data, this supersequence appears to have low potential for generating liquid hydrocarbons (Geoscience Australia, 2010). The final lacustrine sedimentary units belong to the overlying Bronze Whaler Supersequence, which is composed of a dominantly terrestrial lower section and a more marine-influenced upper section (Geoscience Australia, 2010) due to the incursion of marine waters into the evolving seaway between Australia and Antarctica. The lower, non-marine section is deemed to only have moderate propensity for mainly gas and limited oil generation. The more marine-influenced upper sections, by contrast, are regarded as having a much better potential for both oil and gas generation due to higher algal contributions of organic matter. Due to the variation in the relative inputs of marine and non-marine organic matter, it has been speculated that the Bronze Whaler Supersequence would have generated mostly waxy oils and gas, although an increase in marine influence towards the basin depocentre may have resulted in the generation of lighter oil (Geoscience Australia, 2010).

The onset of marine conditions in post-rift sections of the Bight Basin appears to coincide with a first- or second-order rise in global sea level between 134 and 90 Ma, suggesting a possible connection with open-marine conditions along the western Australian margin (Blevin et al., 2000). The rise in global sea level also partially coincides with a period of accelerated subsidence from 97 to 84 Ma, attributed to possible lower crustal processes (Blevin et al., 2000; Totterdell et al., 2000). It is in this setting that the Blue Whale Supersequence, comprising marine siltstone and mudstone, was deposited. This supersequence could exceed 2500 m in thickness in the centre of the Ceduna Sub-basin, with Rock-Eval and TOC data showing good potential for both oil and gas generation (Geoscience Australia, 2010) in samples recovered from the proximal wells in the basin. In the well Platypus-1, both coaly and shaly facies of this supersequence are deemed to have good oil-source potential (Geoscience Australia, 2010).

The White Pointer Supersequence consists of strata deposited in a deltaic environment, comprising fluvial to lagoonal siltstone and claystone. TOC, Hydrogen Index (HI), and Rock-Eval data from several wells indicate that the source potential of this supersequence ranges from good to excellent for both oil and gas. Coaly intervals found in this supersequence are likely to have produced waxy oils (Geoscience Australia, 2010).
The Tiger Supersequence was the final sequence deposited in the transgressive half-cycle and consists of strata deposited in marginal marine to marine settings and reaching a thickness of up to 4000 m in the Ceduna Sub-basin. Currently, only sparse geochemical data exists for this supersequence. Existing TOC and Rock-Eval data from wells drilled through proximal facies indicate poor to fair source potential. However, further basinward, seismic data interpretations infer the presence of thicker shale sequences that may have greater source potential (Geoscience Australia, 2010). Indeed, dredge samples obtained in 2007 from more distal facies of the Eyre Terrace demonstrated the presence of an excellent marine, oil-prone source rock at the base of the Tiger Supersequence (Totterdell et al., 2008; Boreham, 2009).

An important factor to consider during this transgressive half-cycle is that the mid to Late Cretaceous coincides with several global episodes of enhanced organic carbon burial, so called Oceanic Anoxic Events (OAEs; Schlanger & Jenkyns, 1976; Jenkyns, 1986; Arthur et al., 1987; Schlanger et al., 1987). During these short-lived events, elevated primary productivity is generally credited with causing an increased flux of planktonic material to the seafloor, thereby favouring the preservation of this organic matter in areas of oxygen-poor bottom water conditions (Kuyper et al., 2002). OAEs are, among other features, characterised in stratigraphic sections by black shale deposits featuring elevated organic-carbon levels (e.g. Jenkyns & Clayton, 1986). The most significant of all OAEs occurred during the late Cenomanian to early Turonian and is designated as OAE2 (Schlanger et al., 1987). While the majority of records for OAE2 were derived from sediments in the Northern Hemisphere, the occurrence of organic-rich Cenomanian/Turonian sediments at ODP Site 1138 on the Kerguelen Plateau in the southern Indian Ocean (Mohr et al. 2002; Meyers et al. 2007) suggests that mid-Cretaceous anoxic conditions extended to high southern latitudes and that OAE2 affected the entire global ocean (Boreham, 2009). Organic-rich samples recovered from the Eyre Terrace and the Tiger sequence are interpreted to have been related to the OAE2 event due to the presence of dinocyst assemblages composed of morphological variants of Cribroperidinium edwardsii, Cyclonephelium compactum, Eurydinium ingramii and saxoniense, which are associated with the OAE2 in both Hemispheres (Marshall & Batten 1988). In addition, these samples contain strong geochemical signatures of marine anoxia and the presence of age-diagnostic biomarkers (e.g. C26 demethylsterane) consistent with Cenomanian/Turonian (C/T) age (Totterdell et al., 2008).

After the transgressive half-cycle that produced OAE2, a change to regressive environments occurred in the late Santonian between 84 and 65 Ma (Blevin et al., 2000). The start of this half-cycle coincided with the proposed breakup between Australia and Antarctica (Vevers et al., 1991; Sayers et al., 2001, 2003). It was during this time that a large progradational delta, known as the Hammerhead Supersequence, formed. This delta, representing 19 million years of sustained sedimentary deposition, is characterised as sand-rich with strongly progradational stratal geometries (Totterdell & Krasay, 2003).

The Hammerhead Delta is subdivided into three sequence sets, with the lower two (Late Santonian to Campanian) exhibiting strongly progradational geometries, and the upper sequence set (Maastrichtian) aggradational geometries due to, firstly, an increase and, later, a decrease in sediment supply (Totterdell & Krasay, 2003). Core and cuttings and well logs of the supersequence revealed amalgamated sandstone, interbedded claystone-sandstone and massive mudstone units of the Potoroo Formation or its lateral equivalent, the Nurina Formation. In proximal parts of the basin, the Hammerhead Delta consists mainly of amalgamated fluvial channel sandstones, while distal parts consist of basinward-thinning wedges of marine shale at the toe of slope. Between these two end members, 14 other mappable seismic facies have been interpreted to represent coastal and deltaic plain, shallow marine and shelf-margin to slope palaeo-environments. A reduction in
sediment supply at the end of the Cretaceous resulted in the abandonment of deltaic deposition and the onset of a cool-water carbonate margin (Totterdell & Krassay, 2003). Little is known about source rock potential of the deltaic Hammerhead sequence and it remains grossly under sampled. However, it is likely to contain more gas-prone source rocks and coals that are typically associated with deltaic and shallow marine sequences.

Hydrocarbon indications and shows
None of the wells drilled in the Bight Basin have encountered either potential or proven hydrocarbon zones (oil or gas), and hydrocarbon indications and shows are infrequent (Table 1). The only significant oil shows reported are from Greenly-1 (Greenly-1 well completion report), which recovered oil, as a surface scum from a water/oil mixture, and gas (repeat formation test at 4209.2 m). At the time, these represented the first major indications of hydrocarbons in the basin, indicating the presence of an effective source rock, and thereby significantly upgraded the prospectivity of the Ceduna Sub-basin.

Apart from Greenly-1, none of the other wells in the Ceduna and Duntroon sub-basins revealed convincing evidence for hydrocarbon shows. This is particularly relevant for the current phase of exploration in the Bight Basin that centers on exploration permits in the deeper parts of the basin. Only one well, Gnarlyknots-1A, tests this area. While oil, potentially from synthetic-based mud fluorescence, and gas indications were minor, this provided a unique opportunity to test the hydrocarbon generation and migration potential of the area by using fluid inclusions as a tool to access potential hidden oil shows through the GABRP Project 5.3; the results of which are summarised within Box 1.
Table 1: Hydrocarbon indications and shows.

<table>
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<tr>
<th>Well</th>
<th>Year</th>
<th>TD (mRT)</th>
<th>HC shows</th>
<th>Comment</th>
<th>Formation</th>
<th>Interval (mRT)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Borda-1</td>
<td>1993</td>
<td>2800</td>
<td>None reported</td>
<td>–</td>
<td>–</td>
<td>–</td>
</tr>
<tr>
<td>Duntroon-1</td>
<td>1986</td>
<td>3515.6</td>
<td>L1 - oil indication</td>
<td>Fluorescence</td>
<td>Upper Borda</td>
<td>3061; 3200</td>
</tr>
<tr>
<td>Echidna-1</td>
<td>1972</td>
<td>3832</td>
<td>L1 - oil indication</td>
<td>Black soft bituminous material, fluorescence</td>
<td>Lower Borda</td>
<td>2648; 2652</td>
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<tr>
<td>Gnarlyknots-1A</td>
<td>2003</td>
<td>4736</td>
<td>L1 - oil indication*</td>
<td>Fluorescence</td>
<td>Wigunda</td>
<td>4379-4712.5</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>G1 – gas indication</td>
<td>Trace C4</td>
<td>Wigunda</td>
<td>(SWC)</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>L1 - oil indication*</td>
<td>Fluorescence</td>
<td>Wigunda</td>
<td>4505-TD; max</td>
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<td></td>
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<td></td>
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<td></td>
<td>4710</td>
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<td>4685-4725</td>
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<td>Greenly-1</td>
<td>1993</td>
<td>4860</td>
<td>L1, L2 – strong oil</td>
<td>Fluorescence</td>
<td>Wigunda</td>
<td>3430-4542 (SWC,</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>indications</td>
<td></td>
<td></td>
<td>Cuttings)</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>L1 - oil indication</td>
<td>Fluorescence</td>
<td>Platypus</td>
<td>4770-4818</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>L3, G3 – oil/gas show</td>
<td>Recovered oil as surface scum and water/oil</td>
<td>Wigunda</td>
<td>4209.2 (RFT)</td>
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<td></td>
<td></td>
<td></td>
<td>mixture; gas recovered</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Jerboa-1</td>
<td>1980</td>
<td>2538</td>
<td>None reported</td>
<td>–</td>
<td>–</td>
<td>–</td>
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<tr>
<td>Platypus-1</td>
<td>1972</td>
<td>3893</td>
<td>None reported</td>
<td>–</td>
<td>–</td>
<td>–</td>
</tr>
<tr>
<td>Potoroo-1</td>
<td>1975</td>
<td>2924</td>
<td>G1 – gas indication</td>
<td>18,000ppm C1</td>
<td>Platypus</td>
<td>2128-2132</td>
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<tr>
<td>Vivonne-1</td>
<td>1993</td>
<td>3000</td>
<td>None reported</td>
<td>–</td>
<td>–</td>
<td>–</td>
</tr>
</tbody>
</table>

Data summarised from Messent (1998). RFT = repeat formation test. SWC = side-wall core. * Considered likely to be synthetic-based mud fluorescence. L1 = oil indication (fluorescence or cut); L2 = strong oil indication (fluorescence or cut & other oil indication e.g. log anomaly); L3 = oil show (oil recovered from core, test, mud); L4 = potential oil zone (oil show with convincing log anomaly); L5 = proven oil zone (oil flow on test or RFT & log anomaly proving accumulation). G1 = gas indication (anomalously high gas readings); G2 = strong gas indication (anomalously high gas readings and other indications e.g. core, logs or shakers); L3 = gas low on tests; L4 = potential gas zone (gas show with convincing log anomaly or other indication); L5 = proven gas zone (sustained gas flow on test or RFT and log anomaly or pressure data proving an accumulation).
Summary findings of GABRP Project 5.3 fluid Inclusions project

Thirty six samples from 7 exploration wells (Borda-1, Duntroon-1, Greenly-1, Jerboa-1, Platypus-1, Potoroo-1) including the previously untested Gnarlyknots-1A well were analysed using CSIRO’s Grains with Oil Inclusions (GOI™) technique. These analyses identified not only oil-bearing, but in some samples gas-rich, inclusions at low abundance (GOI <0.7%, up to a maximum of 1.1% in Greenly-1) (Kempton et al., 2017). This is positive evidence for widespread oil and gas generation and migration in the Bight Basin. Oil indications are more frequent in intervals from the Late Cretaceous White Pointer, Tiger and Hammerhead supersequences.

Geochemical hydrocarbon fingerprints generated using CSIRO’s Molecular Composition of oil Inclusions (MCI) and gas-isotope techniques were performed on minute quantities of oil and gas extracted from the fluid inclusions. Fluid inclusion (FI) oil from Gnarlyknots-1A (4390–4425 m; Tiger Supersequence) comprises a mixture of types including oil and gas-condensate. The \textit{n}-alkane and biomarker characteristics show a mixed organic matter input from both algae and terrestrial plants and was generated from source rock(s) deposited in suboxic-oxic marine environment(s). The wide range of maturities, 0.65–1.3% VRE, in the FI oil suggests either a mixture of oils generated from different source rocks (Blue Whale and Tiger having potential marine algal inputs, while the White Pointer has potential terrestrial plant input), or from the same source rock at different maturity stages – perhaps an unrecognised paralic facies of the White Pointer, containing both algal and terrestrial organic matter. The gasoline-range hydrocarbons are dominated by toluene and originated, in part, from coeval aqueous inclusions. The bulk $\delta^{13}$C isotopic composition of methane (-28.4 and -28.6 ‰ replicates) and ethane (-17.6 and 18.1 ‰ replicates) indicates a thermogenic origin for these hydrocarbon gases, probably derived from Type III (humic/coaly) organic matter.

By comparison, FI oil from Greenly-1 (4806-4818 m; White Pointer Supersequence) is from oil inclusions only, with no gas-condensate visually detected. The \textit{n}-alkane and biomarker characteristics indicate significant organic matter input from terrestrial plants, and a minor contribution from a lacustrine source, and was generated from a source rock deposited in an oxic, clay-rich fluvio/deltaic depositional environment. This FI oil represents a pristine oil sample that was generated over a narrow maturity range at the early to peak oil window (0.8–1.1% VRE), and lacks the mixed algal input associated with other oil show extracts from the same well. Previous correlations of these oil shows to the Bronze Whaler source sequence are potentially not supported by the lack of algal input to the FI oil, and suggests the White Pointer may be a better source sequence.

To understand the timing of oil migration, the pressure-temperature (P-T) trapping conditions of hydrocarbon inclusions were determined in Gnarlyknots-1A, Greenly-1, Duntroon-1 and Potoroo-1. The intra-Coniacian primary well target in Gnarlyknots-1A (4410-15 m; Tiger Supersequence) was a migration pathway, over an extended period of time, for a variety of hydrocarbon fluid compositions modelled by Petroleum Inclusion Thermodynamic (PIT) as black
oil, light oil, gas-condensate and gas + CO₂. There is good concordance in the measured PT data with independent PT curves from basin models. The earliest oil entrapment took place at a minimum of 58°C, with constrained PT conditions of light oil (240–270 bar; 69.2°C) in the Late Cretaceous at ~70–75 Ma (Campanian). Phase separation of light oil, and entrapment of both gas-rich phases and gas-depleted black oil, occurred at PT conditions (285–308 bar; 80-85°C) in the Late Cretaceous at ~70 Ma (Maastrichtian). Subsequent entrapment of gas-condensate took place at pressure conditions (350–410 bar @ 80°C) reached at ~35–15 Ma (Oligocene to Early Miocene), followed by gas + CO₂ at PT conditions (370–408 bar; 78–88°C) reached at ~27–15 Ma (late Oligocene to Early Miocene). This apparent sequence of hydrocarbon entrapment from oil to gas over time might simply be explained by generation from a single source rock over a range of thermal maturity stages (i.e. paralic facies of the White Pointer?). Late black oil entrapment, at pressure conditions (400 bar) reached at ~18 Ma (Miocene), might be from a different source rock that entered the oil window later in the basin history.

The Cenomanian interval in Greenly-1 (4808–12 m; Tiger Supersequence) trapped remarkably consistent hydrocarbon assemblages modelled by PIT as black oil. While this consistency is also reflected in the measured PT entrapment conditions (270–340 bar; 127–135°C), they are not concordant with the independently modelled PT curve. If the temperatures reflect the entrapment conditions, then this constrains oil charge to the early Miocene (~23–5 Ma) and at a depth equivalent to about 2 km less that modelled. Tertiary carbonate progrades over the Late Cretaceous deltas, and rapid burial peaking around 15 Ma, may go some way to explaining this. More recent charge from the White Pointer source sequence, due to recent sediment loading, is consistent with this timing and is potentially supported by the MCI result. If, however, the pressures reflect the entrapment conditions, then this constrains oil charge in Greenly-1 to the Campanian–early Eocene (~75–52 Ma) and at temperatures about 30°C higher than modelled. One explanation might be a transient period of hotter fluids that were in thermal disequilibrium with the rock.

The Turonian-Santonian interval in Duntroon-1 (2505–10 m; Tiger Supersequence) and the Cenomanian-Santonian interval in Potoroo-1 (1778–86 m; Tiger Supersequence) were migration pathways for a variety of hydrocarbon fluid compositions modelled by PIT as black oil, light oil and, in the case of Duntroon-1, gas (+ N₂ and CO₂). In both these wells located on the margins of the basin, hydrocarbon charge appears to be late. Because pressure evolution over much of the Cenozoic in Duntroon-1 was static, temperature constraints (81–95°C) suggest entrapment from the mid-Miocene (~17 Ma) to recent. The presence of CO₂ and N₂-rich hydrocarbon gases may have volcanic origins similar to those from the Otway Basin. In Potoroo-1, hydrocarbon entrapment occurred either from the Oligocene (~32 Ma), or from the early Miocene (~25 Ma), to recent. Because of its location up-dip of the central Ceduna depocentre, a variety of potential source/charge scenarios could be envisaged including earlier re-migrated oil.
Historical bitumen strandings

In addition to hydrocarbon shows and fluid inclusion studies there has been long standing evidence for bitumen strandings on the beaches of the southern margin of Australia, which could be indicative of active offshore petroleum system in the GAB. Concise summaries of the historical literature on Australian coastal bitumen may be found in Sprigg & Woolley (1963), Padley (1995) and Edwards et al., (2016). The following resumé draws heavily on these sources.

Records of bitumen being washed ashore along the coasts of western and southern Australia date back to the mid-1800s when pioneering sealers and whalers collected this material to caulk their boats (Howchin, 1903; Ward, 1944; Sprigg, 1983). While some of the earliest reports may be found in colonial newspapers, the first official record of coastal bitumen in Australia is contained in a report to the South Australian Surveyor-General. In his capacity as the colony’s Mineral Surveyor, Trewartha (1850) reported finding on the beach at Vivonne Bay, Kangaroo Island, several hundredweight of bitumen in large and small pieces, none of which appeared to be water-worn. The bitumen in question was ‘asphaltum’, later to become known as asphaltite. Tolmer (1882) likewise refers to the ‘asphaltum’ he found on the southern coast of Kangaroo Island in August 1844. Further east, in Victoria, coastal bitumens had been found near Portland and at the mouth of the Bass River in Westernport Bay, some of the latter specimens being noted to “resemble tar” (Brough-Smyth, 1869). ‘Asphaltum’ is the most common variety of Tasmanian coastal bitumen (Hills, 1914; Twelvetrees, 1915; Wade, 1915). The ‘asphaltum’ described by Twelvetrees (1915) appears to physically resemble that which washes ashore in South Australia.

A review of the historical literature (Edwards et al., 2016) revealed that a diverse terminology has been employed to describe what appear to be five distinct varieties of Australian ‘coastal bitumen’:

- **Waxy bitumen**, bitumen, tarballs;
- **Weathered waxy bitumen**, oxidised crude, flaky crude;
- **Asphaltite**, asphaltic bitumen, asphaltum, asphalt, manjak, pitch, mineral pitch, dammar;
- **Oil slick**, oil, oily material, crude, sticky tar, kerosene; and
- **Wax**, ozokerite, paraffin wax, mineral wax, earth wax.

From their descriptions, it is apparent that these stranded bitumens were not of uniform physical appearance, varying from solid to semisolid and liquid states. It is worth noting that the lack of a standard descriptive terminology has, in many instances, made it difficult to decipher exactly what substances were found. Bold type highlights the preferred terminology that will be adopted in this report and subsequent publications.

Early reports of stranded tarballs and ‘asphaltum’ (the two most common varieties) are not restricted to the southern Australian mainland. Other localities include Elcho Island, Northern Territory (Wade, 1924; Hunter, 1964); Mandurah and between Cape Arid and Doubtful Island Bay, Western Australia (Tate, 1883; Sprigg & Woolley, 1963); and the west and south coasts of Tasmania (Hills, 1914; Twelvetrees, 1915, 1917; Wade, 1915; Ward, 1944). Even further afield, similar petroleum substances have been found on the beaches of New Zealand, Chile, South Africa (Ward, 1944) and, more recently, the Seychelles (Plummer, 1996).
Historically, the stranding of coastal bitumen has been reported most frequently at localities in South Australia on the southern Eyre and Yorke Peninsulas, on Kangaroo Island and along the Limestone Coast; and in western Victoria, from Discovery Bay to Cape Otway (Brown, 1898; Ward, 1913; Wade, 1915; Howchin, 1919; Sprigg, 1961, 1962; Sprigg & Woolley, 1963), as highlighted in Figure 6 and Figure 7. Extending from Robe to Portland, the Bonney Coast is arguably the principal locus for bitumen stranding in southern Australia (Figure 7). By comparison, there are notably fewer reports of bitumen stranding along the coastline of the GAB (Figure 6).

These stranded bitumens typically occur on sandy beaches close to the high-water mark, although there are exceptions. Pieces of bitumen and ‘asphaltum’ were found amongst sand hills up to a mile from the coast, particularly near Robe, South Australia (Howchin, 1903; Twelvetrees, 1915; Wade, 1915). A large block of ‘asphaltum’ was encountered 20 feet below the surface, while a well was being sunk adjacent to the Portland-Blowhole road in western Victoria (Reid, 1931).
Understandably, these early reports raised questions regarding the origin of the stranded bitumen, its likely derivation from local or distant oil seeps, and hence its possible implications for the petroleum prospectivity of the sedimentary basin in which the stranding site was located. By the time offshore petroleum exploration commenced in southern Australian waters during the early 1960s, these questions had still not been satisfactorily answered. The uncertain provenance of the bitumen strandings subsequently led to a series of detailed geochemical investigations aimed at establishing how many oil types were represented and their likely source affinities. These investigations focused on material collected from beaches in South Australia and western Victoria (McKirdy & Horvath, 1976; McKirdy et al. 1986, 1994; Padley et al. 1993; Edwards et al. 1998; Hall et al. 2014), Tasmania (Volkman et al. 1992), Western Australia (Currie et al. 1992; Alexander et al. 1994) and the Northern Territory (Summons et al. 1993). With the notable exception of those by Padley/Edwards, these more recent studies typically provided little or no information on the abundance of bitumen found at a given site, its position on the beach, and the shape, size, weight and physical appearance of the individual specimens collected.

In 1983 the South Australian Department of Mines and Energy (SADME) conducted a major survey of stranded bitumen along the South Australian and western Victorian coastline, involving the collection of new samples for geochemical analysis (McKirdy, 1984a, b), with a view to promoting petroleum exploration within the adjacent offshore Otway and Duntroon Basins. The same agency sponsored follow-up beach surveys during 1990–1992 as part of a PhD project (Padley, 1990, 1992, 1995).
latter study was the first to yield a quantitative estimate of the natural hydrocarbon loading of Australian ocean beaches (as documented in Edwards et al., 2016).

Evidence for seeps
To attempt to identify the source of some of the asphaltites and help de-risk offshore hydrocarbon exploration, several seepage studies have been conducted within the GAB, primarily focusing on the Ceduna sub-basin. These have included remote sensing synthetic aperture radar (SAR) data collection, seismic interpretations and vessel-based investigations.

Remote sensing and seismic studies
Remote sensing studies were undertaken in the GAB by Australian Geological Survey Organisation (AGSO, now GA) both in 1999 (Stuckmeyer, 2000) and 2005/2006.

These combined studies collected 201 synthetic aperture radar (SAR) captures over the wider GAB, using a combination of European remote sensing satellite (ERS) and RadatSat SAR satellite platforms, to evaluate hydrocarbon migration and seepage. Within these captures 183 SAR anomalies were identified, none of which were interpreted as high confidence slicks (ranked 2 and below). Seventeen slicks were identified as rank 2 slicks with moderate confidence, and 97 as rank 3 slicks (lowest rank) with moderate to low confidence. All other SAR anomalies were assigned to remnant pollution or natural films (Figure 8).

These data were compared with legacy airborne laser fluorosensor (ALF) and 2D seismic data to provide an assessment of the likely charge characteristics of the region. The results of the study suggested that there was active liquid seepage with the Great Australian Bight basin system albeit restricted. Direct correlation of the individual seeps with geological features was not possible as none of the reported seep locations coincided with subsurface fluid escape features identified from the limited seismic data available in the basin at the time.

The high proportion and distribution of the SAR anomalies identified in the west of the Ceduna sub-basin in EPP 43 during the 2005/2006 capture campaign was distinctly different from those encountered in the 1999 acquisition program data. In part these data were used in the planning of the seeps directed activities of the R/V Southern Surveyor Survey SS01/2007.

These remote sensing studies have highlighted that slicks were more abundant across the NW Shelf than in the GAB. This may be due to a number of reasons:

- Differences in water depth, swell height and water temperature – an example being consistently high swells that may account for the small size and scattered nature of slicks;

- High seal and trap integrity in the greater part of the GAB basin system;

- Low number of seeps could reflect a low density of hydrocarbon accumulations; and

- Episodic hydrocarbon leakage (Struckmeyer, 2000)
Figure 8: Combined historical and capture program (SAR Slick Features) sea surface synthetic aperture radar anomalies for the Great Australian Bight, plus seismic indications of fluid escape in the Duntroon sub-basin.

Vessel-based investigations

There have been eight known prior marine research surveys undertaken by Geoscience Australia and CSIRO studying aspects of the geology of the GAB between 1989 and 2013. Whilst these surveys have collected numerous samples across the bight (Figure 9) where samples have been geochemically characterized none have identified thermogenic hydrocarbons indicative of deep source seepage.

Only two of the eight research surveys in the GAB have specifically attempted to delineate hydrocarbon seepage (viz. SS01/2007; Totterdell & Mitchell 2009; ss2013_C02, Williams et al., 2013a,b). Whilst the number of marine surveys performed in the GAB appears to be numerous, only recent advances in marine survey technologies, and the collection of new seismic and remote sensing data sets have permitted a more detailed understanding of the GAB geology and planning of a more precise survey.

The ss2013_C02 voyage was primarily focused on benthic and pelagic ecology studies with only limited acoustic reconnaissance opportunities undertaken for the delimitation of potential seepage (using sub-bottom profiler, single beam echo sounder water column acoustics and multibeam echo sounder swath mapping (Williams et al., 2013b)). These operations predominantly occurred at the transition between deepwater to abyssal slope between 131°5'-133°E and 35°-36°S. The sites investigated were selected on the basis of 2D seismic data and synthetic aperture radar captures.
concurrent with the voyage, which identified potential sea surface slicks (see Box 2, Figure 8). Several low rank acoustic anomalies were identified during this voyage within both the sub-bottom profiler data and the single beam echo sounder data, although these data are equivocal (Ross et al., 2017).

![Combined historic seafloor sampling sites from surveys undertaken in the Great Australian Bight.](image)

The ss201/2007 multifaceted Geoscience Australia led survey (Totterdell & Mitchell, 2009) aimed to investigate potential natural hydrocarbon seepage at sites across the Ceduna sub-basin, which could provide evidence for the presence of active petroleum systems. The survey investigated six potential seepage locations (Areas 1-3, 7-9, Figure 10) selected on the basis of geophysical and remote sensing seepage indicators (Totterdell & Mitchell, 2009).
Figure 10: Bight Basin Sampling and Seepage Survey SS01/2007: survey areas and location of sampling sites (Totterdell & Mitchell, 2009).

These sites were initially characterised using acoustic methods, namely a hull-mounted single dual-frequency echo sounder for the detection of water column hydro-acoustic flares and a multibeam echo sounder system for bathymetric mapping. Shallow subsurface geophysical data were also collected through the use of a sub-bottom profiler system. Detailed site characterisation was achieved using a towed video camera system and side-scan sonar system (Area 1) only. Sediment samples were collected with a combination of grab, dredge and coring techniques for detailed geochemical investigation, including headspace gas and high-molecular-weight hydrocarbon analysis.

Data collected within Area 1 and on the transit to Area 2 delineated a number of potential seep-related features, including areas of high seabed reflectivity observed in side-scan sonar data, potential hydro-acoustic flares and amplitude anomalies within sub-bottom profiler data. These features were observed in the vicinity of subsurface faulting and SAR slick anomalies. However, sampling within these areas failed to identify thermogenic hydrocarbons in both the headspace gas and the high-molecular-weight hydrocarbons.

Multibeam swath data collected within Area 3 allowed mapping of reactivated underlying basin margin faults which form seafloor scarps. An area of seabed with circular depressions was identified on a plateau in the northern part of Area 3 at a water depth of about 700 m. These depressions, ~150 m in diameter and 5 m deep, are located on a ~5° SE slope. These features were interpreted by Totterdell & Mitchell (2009) to be possible pockmark features related to fluid escape that is currently active, or alternatively recent palaeo-fluid escape features. Once again sampling of these features and geochemical analysis of the samples failed to positively detect thermogenic hydrocarbons.
The remaining areas studied during the SS01/2007 voyage failed to unequivocally identify seepage and Totterdell & Mitchell (2009) cite several possible reasons for this, including hydrocarbons not being preserved in the very shallow section (<10 m) that was sampled by the gravity corer, poor targeting of potential seepage sites, or the lack of any hydrocarbon seepage. In particular they advocate for the use of equipment capable of taking deep cores in order to access deeper intervals where hydrocarbons may be present.
Box 2

**Seepage indicators from GABRP Project 5.1 seeps and leakage project.**

To identify potential sites of hydrocarbon seepage a multidisciplinary approach was taken. Of the 81 areas of interest identified from the rapid screening of 2D seismic data for leakage indicators within the Ceduna sub-basin (Figure 8) no one area displayed unequivocal evidence of fluid leakage through the subsurface to the seafloor (Ross et al., 2017).

Sea-surface slick studies included the acquisition of four additional SAR scenes which augmented the 231 publically available scenes. Twenty three sea-surface anomalies were identified within the four additional scene captures which clustered in proximity to anomalies identified in prior scene captures. These data were used, in conjunction with the areas of interested identified for 2D seismic data screening, to determine areas for subsequent reconnaissance on the SS2013_C02 voyage.

Of the areas studied during the SS2013_C02 voyage and the subsequent IN2015_C02 voyage, there is only weak evidence for possible seepage from the data collected, in part due to the very limited time spent on these activities as part of the voyages. This outcome is not a definitive indication of the absence of seepage in the basin, as the basin comprises a very large area.

Of the 84 samples that were collected during the SS2013_C02 voyage (Figure 8) and by a 2013 BP chartered *M/V Fugro Southern Supporter* voyage, quantitative analysis of seawater extracts indicated that polycyclic aromatic hydrocarbons (PAHs) are largely absent in and around the BP permits in the GAB. Naphthalene, a highly volatile bicyclic aromatic hydrocarbon, detected in small amounts in the majority of the samples, is probably a contaminant cumulatively inherited from solvents, glassware and/or unknown source(s).

The qualitative organic geochemical analyses of extractable organic matter isolated from seawater, seabed sediment, and GO net samples, and headspace gases extracted from multicore/piston core samples collected from the GAB region indicate that petroleum hydrocarbons are either absent or below the limits of detection, pointing to an absence of active hydrocarbon seepage in and around the sampling region. The trace amounts of thermogenic hydrocarbons found in two water samples and a GO net sample are likely to be inherited from a contaminant oil of an unknown source.

The seabed sediment samples contain variable amounts of extractable organic matter (0–3200 ppm) in which there are no detectable petroleum hydrocarbons (Ross et al., 2017).

**Formation of coastal bitumen and potentially analogous systems**

As no seafloor seeps for the asphaltites and waxy bitumen collected along the South Australian coastline have yet been explicitly located or characterised, the exact processes responsible for their
formation cannot be easily determined. This section provides a brief summary of the potential origin mechanisms for the Southern Australian margin asphaltites and waxy bitumen (tarballs), followed by overviews of two of the best-studied potential analogues: the Gulf of Mexico and the Santa Barbara Channel.

Proposed origin processes for South Australian asphaltites and waxy bitumen

Proposed origin of waxy bitumen
The origin of the waxy bitumen has largely been ascribed to unidentified seafloor tar seeps in the Indonesian Archipelago (Edwards et al., 2016, 2017). These bitumens, which are positively buoyant, become entrained in the ocean currents of the Indonesian Throughflow and are ultimately transported to Australia’s southern margin by the Leeuwin Current (Figure 11). As the documentation of waxy bitumen strandings predates significant petroleum exploration in Indonesia and crude oil importation by tanker, the waxy bitumens are unlikely to be the result of accidental spillage (Edwards et al., 2016). However, given the presence of multiple waxy bitumen families (McKirdy et al. 1984a, b; Padley, 1995; Edwards et al., 2016, 2017), previous studies also note the possibility that some varieties of waxy bitumen, specifically those lacking botryococcane, may be derived from sedimentary basins along Australia’s southern margin (McKirdy et al., 1994; Padley et al., 1995).

Proposed origin of asphaltites
Presently two competing hypotheses exist to explain the origin of the asphaltites: (1) degradation of a natural oil slick on the sea surface which progressively degrades into a tar ‘mousse’ before completely solidifying (Edwards et al., 1998; Logan et al., 2010) and (2) seepage of asphalt directly onto the seafloor (Hall et al., 2014). Previously the asphaltites were also hypothesised to be anthropogenic in origin, as many identified stranding sites coincided with historical whaling stations (Edwards et al., 1998). This suggested the asphaltites may have been caulking bitumen stored in casks buried on the whaling station beaches which were eroded and delivered to the ocean through storm activity. The hypothesis was similarly supported by historical accounts of the asphaltites being used to seal wooden ships. However, this interpretation has since been falsified by geochemical comparison of the asphaltites with caulking tars recovered from historical shipwrecks which revealed no relationship between them (Smart, 1999). Additionally, the continued stranding of fresh specimens suggests that their source remains active.

The degradation of a surficial oil slick to form a semi-solid bitumen was proposed by both Edwards et al. (1998) and Logan et al. (2010). This process has been documented in the Santa Barbara Channel (see analogues section below) where tar stranding is substantially greater in the summer months due to calmer sea-surface conditions which allow tar ball formation. In winter, due to wave action and winds from storms, the oil slicks break apart and are dispersed before they can form tar balls and be deposited on the shoreline (Lorenson et al., 2009). The stranding of asphaltite along the Limestone Coast also occurs predominantly during the winter months due to a combination of increased storm frequency and the dominance of the Coastal Current along Australia's southern margin (Padley, 1995; Edwards et al., 2016).